Practical Multigrid Methods for Momentum Balance in Ice Sheets

This talk: http://59A2.org/files/20150202-LIWGMultigrid.pdf

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Why do we need scalable solvers?

Increasing resolution

- larger problem sizes
- more 3D effects visible
- more time steps \Rightarrow smaller budget per time step
- Sequence of simulations data assimilation, UQ
- All other costs typically linear in problem size

Review: two definitions of scalability

"Strong scaling"

- execution time (7) decreases in inverse proportion to the number of processors (p)
- fixed size problem (N) overall
- often instead graphed as reciprocal, "speedup"
- "Weak scaling" (memory bound)
 - execution time remains constant, as problem size and processor number are increased in proportion
 - fixed size problem per processor
 - also known as "Gustafson scaling"



The easiest way to make software scalable is to make it sequentially inefficient. – Gropp 1999

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Is multigrid needed?

Long-range coupling is slow to converge using local methods

- Iteration count proportional to diameter of support of Green's functions
- How local are the Green's functions?

Columns of the inverse matrix

$$u(x) = \int_{y \in \Omega} G(x, y) f(y)$$

- Sticky, flat bad Green's functions local (and SIA is accurate)
- Slippery bed (ice shelf), steep topography at high resolution
- Pressure: surface is Dirichlet boundary condition, causes rapid decay



Bathymetry and stickyness distribution

- Bathymetry:
 - Aspect ratio $\varepsilon = [H]/[x] \ll 1$
 - Need surface and bed slopes to be small
- Stickyness distribution:
 - Limiting cases of plug flow versus vertical shear
 - Stress ratio: $\lambda = [\tau_{xz}]/[\tau_{membrane}]$
 - Discontinuous: frozen to slippery transition at ice stream margins
- Standard approach in glaciology:

Taylor expand in ε and sometimes $\lambda,$ drop higher order terms.

- $\lambda \gg 1~$ Shallow Ice Approximation (SIA), no horizontal coupling
- $\lambda \ll 1~$ Shallow Shelf Approximation (SSA), 2D elliptic solve in map-plane
 - Hydrostatic and various hybrids, 2D or 3D elliptic solves
- Bed slope is discontinuous and of order 1.
 - Taylor expansions no longer valid
 - Numerics require high resolution, subgrid parametrization, short time steps
 - Still get low quality results in the regions of most interest.
- Basal sliding parameters are discontinuous.

Hydrostatic equations for ice sheet flow

- Valid when $w_x \ll u_z$, independent of basal friction (Schoof&Hindmarsh 2010)
- Eliminate *p* and *w* from Stokes by incompressibility: 3D elliptic system for u = (u, v)

$$-\nabla \cdot \left[\eta \begin{pmatrix} 4u_x + 2v_y & u_y + v_x & u_z \\ u_y + v_x & 2u_x + 4v_y & v_z \end{pmatrix}\right] + \rho g \overline{\nabla} h = 0$$

$$\eta(\theta, \gamma) = \frac{B(\theta)}{2} (\gamma_0 + \gamma)^{\frac{1-\pi}{2n}}, \qquad \mathfrak{n} \approx 3$$
$$\gamma = u_x^2 + v_y^2 + u_x v_y + \frac{1}{4} (u_y + v_x)^2 + \frac{1}{4} u_z^2 + \frac{1}{4} v_z^2$$

and slip boundary $\sigma \cdot n = \beta^2 u$ where

$$\begin{split} \beta^2(\gamma_b) &= \beta_0^2 (\varepsilon_b^2 + \gamma_b)^{\frac{\mathfrak{m}-1}{2}}, \qquad 0 < \mathfrak{m} \leq 1\\ \gamma_b &= \frac{1}{2} (u^2 + v^2) \end{split}$$

 Q1 FEM with Newton-Krylov-Multigrid solver in PETSc: src/snes/examples/tutorials/ex48.c



- Bathymetry is essentially discontinuous on any grid
- Shallow ice approximation produces oscillatory solutions
- Nonlinear and linear solvers have major problems or fail
- Grid sequenced Newton-Krylov multigrid works
 - as well as in the smooth case



Figure: Grid sequenced Newton-Krylov convergence for test *Y*. The "cliff" has 58° angle in the red line (12×125 meter elements), 73° for the cyan line (6×62 meter elements).

Strong scaling on Blue Gene/P (Shaheen)



Figure: Strong scaling on Shaheen for different size coarse levels problems and different coarse level solvers. The straight lines on the strong scaling plot have slope -1 which is optimal.

Weak scaling on Blue Gene/P (Shaheen)



Figure: Weak scaling on Shaheen with a breakdown of time spent in different phases of the solution process. Times are for the full grid-sequenced problem, not just the finest level solve.

One high-accuracy solve costs 30 times as much as a residual evaluation

about 15 to reach truncation error

1000 times faster than some popular methods e.g. Lemieux, Price, Evans, Knoll, Salinger, Holland, Payne 201 (J. Computational Physics) — Actual speedup subject to Amdahl's Law

(Brown, Smith, Ahmadia 2013, SIAM J. Scientific Computing)

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Algebraic multigrid for Hydrostatic

- Easy to use: assemble a matrix and throw it over the wall
- Higher setup costs, lower arithmetic intensity
- AMG uses heuristics to diagnose anisotropy; varies by discretization
- Need to represent rotational modes
 - Smoothed aggregation takes a "near null space" (translation plus rotation)



Eigen-analysis plugin for solver design

Hydrostatic ice flow (nonlinear rheology and slip conditions)

$$-\nabla \left[\eta \begin{pmatrix} 4u_x + 2v_y & u_y + v_x & u_z \\ u_y + v_x & 2u_x + 4v_y & v_z \end{pmatrix} \right] + \rho g \nabla s = 0, \quad (1)$$

- Many solvers converge easily with no-slip/frozen bed, more difficult for slippery bed (ISMIP HOM test C)
- Geometric MG is good: $\lambda \in [0.805, 1]$ (SISC 2013)



HPGMG-FE https://hpgmg.org



Δ

Conservative (non-Boussinesq) two-phase ice flow

Find momentum density ρu , pressure p, and total energy density E:

$$(\rho u)_t + \nabla \cdot (\rho u \otimes u - \eta Du_i + \rho 1) - \rho g = 0$$

$$\rho_t + \nabla \cdot \rho u = 0$$

$$E_t + \nabla \cdot ((E + \rho)u - k_T \nabla T - k_\omega \nabla \omega) - \eta Du_i : Du_i - \rho u \cdot g = 0$$

- Solve for density ρ, ice velocity u_i, temperature T, and melt fraction ω using constitutive relations.
- This and many other formulations lead to a Stokes problem



The Great Solver Schism: Monolithic or Split?

Monolithic

- Direct solvers
- Coupled Schwarz
- Coupled Neumann-Neumann (need unassembled matrices)
- Coupled multigrid
- X Need to understand local spectral and compatibility properties of the coupled system

Split

- Physics-split Schwarz (based on relaxation)
- Physics-split Schur (based on factorization)
 - approximate commutators SIMPLE, PCD, LSC
 - segregated smoothers
 - Augmented Lagrangian
 - "parabolization" for stiff waves
- X Need to understand global coupling strengths
- Preferred data structures depend on which method is used.
- Interplay with geometric multigrid.

Stokes

Weak form of the Newton step

Find (u, p) such that

$$\int_{\Omega} (Dv)^{T} [\eta 1 + \eta' Dw \otimes Dw] Du$$
$$-p \nabla \cdot v - q \nabla \cdot u = -v \cdot F(w) \qquad \forall (v,q)$$

Matrix

$$\begin{bmatrix} \mathbf{A}(\mathbf{w}) & \mathbf{B}^T \\ \mathbf{B} \end{bmatrix} \begin{pmatrix} u \\ p \end{pmatrix} = - \begin{pmatrix} F_u(\mathbf{w}) \\ 0 \end{pmatrix}$$

Block factorization

$$\begin{bmatrix} A & B^T \\ B \end{bmatrix} = \begin{bmatrix} 1 \\ BA^{-1} & 1 \end{bmatrix} \begin{bmatrix} A & B^T \\ S \end{bmatrix} = \begin{bmatrix} A \\ B & S \end{bmatrix} \begin{bmatrix} 1 & A^{-1}B^T \\ 1 \end{bmatrix}$$

where the Schur complement is

$$S = -BA^{-1}B^T.$$



Hardware Arithmetic Intensity

Operation	Arithmetic Intensity (flops/B)
Sparse matrix-vector product	1/6
Dense matrix-vector product	1/4
Unassembled matrix-vector product, residual	\gtrsim 8

Processor	STREAM Triad (GB/s)	Peak (GF/s)	Balance (F/B)
E5-2680 8-core	38	173	4.5
E5-2695v2 12-core	45	230	5.2
E5-2699v3 18-core	60	660	11
Blue Gene/Q node	29.3	205	7
Kepler K20Xm	160	1310	8.2
Xeon Phi SE10P	161	1060	6.6
KNL (DRAM)	100	3000	30
KNL (MCDRAM)	500	3000	6

Outlook

- Choose suitable technology
- Geometric multigrid is simple and has low setup cost
- Algebraic multigrid has higher setup, more finicky to discover anisotropy
- Stokes problems
 - block factorization is easiest (all run-time options in PETSc)
 - coupled MG is worth considering
- Newton linearization of sliding
- Mind the external factors